

Understanding DPR DSD Parameters on a Global Scale

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I. INTRODUCTION

The Global Precipitation Measurement (GPM) core satellite carries the first spaceborne dual-frequency precipitation radar (DPR) at Ka (35.5 GHz) and Ku (13.6 GHz) frequencies. One of the advancements is that it quantitatively estimates the precipitation particle size distributions. The estimated size parameter, often called the "mass weighted mean diameter", Dm, offers new physical insights into microphysical properties of precipitation around the globe.

The Level-3 DPR product provides gridded precipitation quantities and statistics based on the Level-2 swath data. We use 4 years of DPR L-3 data with 5° grid resolution to quantify the horizontal and vertical distribution of Dm, along with the corresponding Nw and precipitation rates. We also investigate its seasonal variations and relationships with surface and rain types.

II. ANNUAL MEAN OF PRECIPITATION SIZE, NUMBER, AND RATE

Based on Seto et al. (2013) and Williams et al (2014), the drop size distribution (DSD) in the DPR algorithm uses a normalized gamma distribution function, where the number density is:

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$$N(D) = N_w f(D; D_m) \ , \qquad N_w = \frac{4^4}{\pi \rho_w} \left(\frac{a}{D_m^4}\right) \ \text{and} \qquad D_m = \frac{\sum_{D_m \text{tot}}^{D_m \text{tot}} D_d^4 dD}{\sum_{D_m \text{tot}}^{D_m \text{tot}} D_d^3 D}$$

The precipitation rate is:

$$R = N_w H(D_m)$$
, where $H(D_m) = 0.6\pi \times 10^{-3} \int_{D=0}^{\infty} V(D) D^3 f(D; D_m) dD$

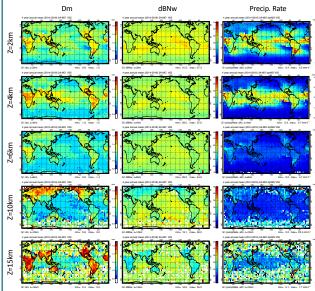
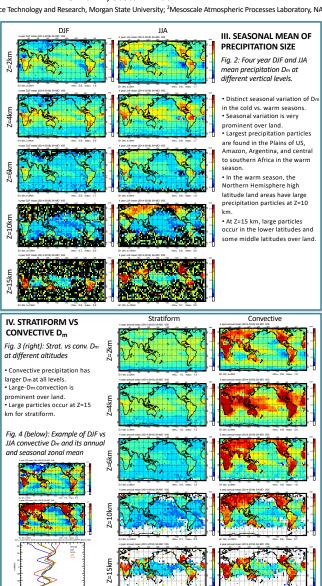


Fig. 1: Four year annual mean DSD parameters and precipitation rate at different vertical levels.

- Dm is larger over land than over the ocean.
- Shallow cumulus regions in the eastern side of ocean basins have smaller particle sizes than the western ocean basins.
- Arid and desert regions over land have small particle sizes.
- In the middle to high latitudes of Eurasian and North America, the particle sizes are large at Z=10 km.
- At Z=15 km, large particles are present in the lower latitudes over land.

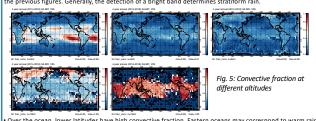
ieto, S., T. Iguchi, and T. Oki, 2013: The basic performance of a precipitation retrieval algorithm for the Global Precipitation Measurement Mission's single/dual-freque adar measurements. IEEE Trans. Geosci. Remote Sens., 51, 5239-5251, doi:10.1109/TGRS.2012.2231686.

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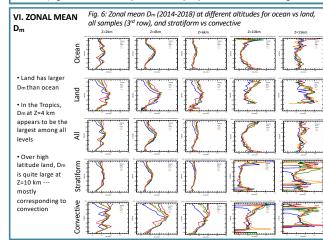
V. CONVECTIVE FRACTION

The fraction of convective precipitation is calculated to help us understand the characteristics of Dm shown ir the previous figures. Generally, the detection of a bright band determines stratiform rain



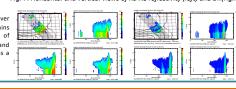
Over the ocean, lower latitudes have high convective fraction. Fastern oceans may correspond to warm rain At Z=10 km, high convective fraction in the Northern Hemisphere high latitudes corresponds to large Dm.

At 7=15 km, high convective fraction may indicate that only deep convection can reach such high altitudes



VII. A SQUALL LINE Fig. 7: Horizontal and vertical views of Ku NS reflectivity (left) and Dm(right) CASE

Example of squall line over the Southern Great Plains showing the variability of Dm in convective and stratiform regions and as a function of altitude.



VIII. SUMMARY

We analyzed four years of GPM DPR data to study precipitation particles sizes over the globe. Particle size distributions show clear contrasts over land vs. ocean, with distinct variations as a function of season, latitude, and altitude. Large particles are generally associated with convective precipitation, except for warm rain regions over the low latitudes of eastern ocean basins. Large particles are found in northern Eurasia and North America at high altitudes in the warm season, which corresponds to large convective fractions reaching that level. The particle sizes of stratiform precipitation at Z=15 km are large, which will be further investigated. Individual cases are being selected to improve the understanding of the statistics of Dm.